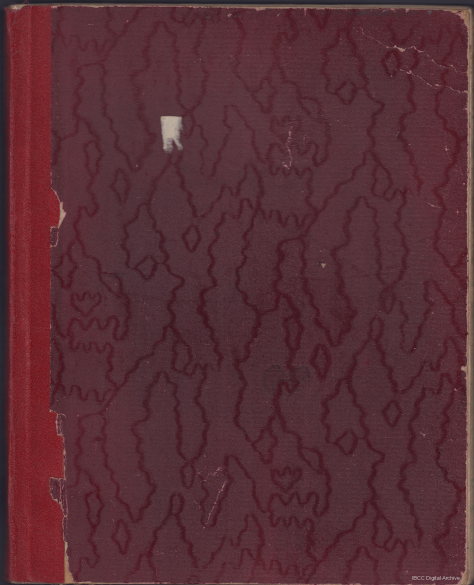




METEOROLOGY

Index

1. Introduction
2. The Atmosphere.
 3. Pressure.
 - a) Pressure Distribution.
 - b) Pressure and Temperature.
 - c) The Inversion.
 - d) The Adiabatic Lapse Rate.
 - e) Conditions for Stability in the Atmosphere.
 4. Barometers - Altimeters.
3. Wind - Measuring & Estimating Wind Speed & Direction
4. Clouds.
 - a) Humidity & Dew Point.
 - b) Wet and Dry Thermometer.
 - c) Formation of Clouds.
 - d) Types of Clouds - Index.
5. Visibility in Clouds. Fog. Mist. Haze etc.
6. Precipitation from Clouds. Rain, Snow, Sleet, Hail, etc.
7. Sea weather, and Thunderstorms.
8. RAF Meteorological Service, & Basis of Forecasting.
 1. Factors affecting the Weather.
 2. Polar & Tropical Air Weather.
9. Fronts, Cold and Warm.
10. a. The Beaufort Weather Notation.
11. b. Information on a Weather Map.

1. Introduction.

Officers need to follow Meteorology for the following reasons.

- a) To their efficiency and safety.
- b) To understand the behaviour of the atmosphere.
- c) To understand the weather forecast and make
- d) To bring back accurate and useful information and therefore assist in the planning of future Operations.

2. The Atmosphere.

As well as Oxygen and Nitrogen, the atmosphere consists of Water Vapour, Water Droplets and Solid Particles.

3. Pressure.

Pressure is the weight of air per unit area, pressing down on that point.

Pressure always decreases - as we ascend. It is measured in millibars - 1 millibar is approx. 1 gram on 1 square centimeter. Pressure decreases with height at the rate of about 1 millibar every 30 ft near the ground, gradually becoming 1 millibar per 60 ft at about 20,000 ft. The average M. S. L. pressure - 1013.2 mb.

a) Type of Pressure Distribution

i) Depression or Low

→ Wind at 15,000 ft.

→ Wind at Surface



Winds at 15,000 ft. follows the Isobars and blow in an anticlockwise direction around the centre of Low.

The speed of the wind is estimated by the closeness of the Isobars.

Due to friction and the uneven surface of the earth the wind at the surface will move round the centre at an angle of about $30-50$ degrees to the Isobars. The surface wind blow rugged to the low pressure at a speed of about $\frac{1}{2}$ that of the wind at 15,000 + 20,000 ft.

Over the sea the angle is $15-16^\circ$ and the speed is about $\frac{3}{4}$ of the 2,000 ft wind. Therefore when we ascend from the ground up to 2,000 ft the wind will always VEER and increase.

There are 3 kinds of Depression

- Deep Depression - one which has a barometric reading at the centre of the Low of 980-970 etc.
- Intense Depression - one which has Isobars very close together and therefore has very strong winds.
- Shallow Depressions - one which has Isobars a long way apart and therefore relatively light winds. Bad weather system, plenty of cloud and precipitation.

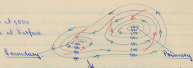
Kind of Weather from a Depression.

Bad weather system, plenty of cloud and precipitation. The depressions can move in any direction and at any speed. More than 50% of the Depressions in this part of the world travel from S.W. across the British Isles towards Russia. The speed varies from 5 to 30 mph - some may travel up to 50 mph and some have been known to be almost stationary.

2) Secondary Depression

Has the same properties as the ordinary Depression. P.T.O.

- Wind at 1500
→ Wind at Surface



A Secondary Depression is a separate area of low pressure within the Primary. It tends to rotate round the Primary in an anti-clockwise direction like the system as a whole moves forward.

The Primary or Old Depression may be moving slowly whilst the Secondary or new Depression is moving quicker. The secondary may have a lower or deeper pressure in the centre than the Primary, also it may be more intense.

Weather with a Secondary Depression is similar to the Primary but the winds - due to the Secondary are usually much stronger.

2) Trough of Low Pressure

The main system is a low pressure system but instead of more or less

circular isobars, you get a nose of isobars being developed.

Sometimes the nose develops and forms a V shaped trough in which case it is called a V shaped Depression.

Trough of Low Pressure



- Wind at 1500
→ Wind at Surface



The worst weather is experienced along the line of the trough or V.

3) High Pressure - Anticyclone

- Wind at 1500
Wind at Surface



The isobars are more or less concentric curves, winds are light. The highest pressure is at the center and the isobars are not so close together as in a Depression. The wind at 1500-2000 feet follows the isobars flowing in a clockwise

direction round the centre of High.

At surface level the winds blow outwards at an angle of $20-30$ degrees to the Isobars.

It indicates good weather and although it can move in any direction with a speed of usually less than 3 m.p.h., it can stay put in any one region for a long time.

A system of light winds (Isobars far apart).

The winds at $1500-2000$ ft are in the neighbourhood of less than 10 to 15 m.p.h.

Two main sorts of weather.

1. Fine weather.

2. Anticyclonic gloom - may be caused by

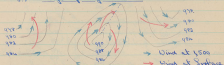
a) % of Stratus or Strato-cumulus clouds, or

b) Dg Radiation fog.

In Winter. Time a) + b) will persist throughout the Day and Night.

In Summer. Time a) + b) will form at night and disappear in the daytime. Owing to very light winds, these conditions will not be swept away but will linger on.

5.1 Ridge of High Pressure



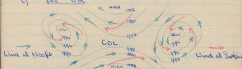
A Ridge of High Pressure is a bulge towards the Pole in a High Pressure System usually situated between two Low Pressures.

The wind at $1500-2000$ is clockwise round High and round the two Low Pressures in an anticlockwise direction and given in the direction of the Isobars.

A fine weather system but differs slightly from that of the High Pressure System itself in the time it persists.

On each side of the wedge there is likely to be a Depression moving South Eastwards at 20 m.p.h. Therefore the wedge is likely to move in the same direction and at the same speed therefore the fine weather of the wedge cannot last very long.

a) The Col.



The Col is the central area made up by 2 Low Pressure areas and 2 High Pressure areas. No barbes exist.

The winds are therefore are indefinite and very light. There may even be a dead calm.

Two main sorts of weather from a Col.

- 1) Winter Time. - Fog or overcast sky due to the mixing of air.
- 2) Summer Time. - Thunderstorms due to the cumulo-nimbus clouds which develop over the area due to the excessive heating of the Earth's surface beneath it, and as there is no cooling wind thunderstorms will develop. Not likely to last for more than 12 hours.

b) Pressure and Temperature

Is measured in degrees Centigrade or Fahrenheit. The volume of the air depends upon its absolute temperature which equals Centigrade plus -273° .

Horizontally temperature from the Equator to the Poles, giving Temperature Zones.

1. 25° N to 5° S of the Equator is called Tropical.
2. 25° - 40° N to 5° S of the Equator. Temperate.
3. 40° - Pole N to S. Polar.

5 Zones

Vertically, temperature usually decreases, the average rate being 3°F or 1.9°C per thousand feet, but temperature ceases to decrease at the Tropopause, the height of which varies in this country from 23,000 ft in the winter to 43,000 feet in the summer. The mean height of the Tropopause over the Poles is 6 1/2 miles and over the Equator 11 miles.

The Density decreases as we ascend because of the decrease of pressure and in spite of the decrease of temperature.

The effects on an aircraft are

- a) Decrease Lift
- b) Decrease in Drag
- c) Decrease Thrust
- d) Decrease in Engine Power (Lack of Oxygen)
- e) Decrease of Indicated Air Speed.

Therefore the A/c has a 'ceiling' and also optimum height. (the height of maximum performance)

The main source of atmospheric heat

Heat is obtained from the Sun by Radiation. This radiated heat warms the Earth to varying degrees, dependent upon the specific heat of the substances of which the Earth is composed. (Some is not absorbed but is reflected back into space.)

The layer of air next to the surface of the warmed Earth is heated by conduction, this warmed air rises to the upper layers of the atmosphere by convection.

The Relation Between Temperature and Height

Normally temperature diminishes at height above the ground. The rate of 'fall off' is called the Lapse Rate

Normal Lapse Rate is 3°C every 1000 ft, but the actual temperature Lapse Rate may differ considerably from this average.

c). The Inversion

Occasionally temperature increases with height over a short range especially near the ground, this is called an Inversion.

Near the ground an Inversion is not normally produced during a clear Winter night, when the Earth radiates its heat into space and hence becomes cooled, cooling the air in immediate contact with it to such an extent that the air at some distance above the Earth is warmer than it. An Inversion is then established.

d). The Adiabatic Lapse Rate

When a mass of air is heated, it expands and rises and arrives at an area of lower Pressure and as a result it expands. Due to this expansion the temperature drops but there is no loss of temperature to the outside surroundings.

The fall in temperature is due to the

internal expansion of the mass of air and is called adiabatic cooling.

The Dry & Adiabatic Surface Rate is 5.4°F per 1,000 feet.

The Wet Adiabatic Surface Rate

On cooling, wet & saturated air, some moisture condenses into Water Droplets. The latent heat thus released stops & prevents the air from cooling off at the same rate as previously. This is the Wet Adiabatic Surface Rate which is 2.9°F per 1,000 ft.

c) Conditions for Stability in the Atmosphere

a) Atmospheric Stability

When the normal Surface Rate for the day is less than the Dry Adiabatic Surface Rate, then the atmosphere is said to be stable i.e. rising air is not going to be continually rising and when a rising mass of air attains the same temperature of the surrounding air, the mass of air tends to be thrown back.

b) Atmospheric Unstability

When the normal Surface Rate for the day is very steep, such as is obtained

during a very hot day and greater than the Dry Adiabatic Surface Rate, then the atmosphere is said to be unstable, and air will continue to rise and Thunderstorms are likely to appear.

Atmospheric Stability

Atmospheric Unstability

4000ft		45°	45°	
3000ft		51°	55°	
2000ft	(57°)	57°	60°	(74°)
1000ft	(63°)	63°	72°	(82.5°)
0 ft.	(69°)	69°	80°	(90°)

10. Barometers and Altimeters

- a) The Mercury Barometer - is used to measure pressure.

The air pressure balances a column of Mercury and the height of it is measured in millibars and is corrected for.

- a) Temperature - Standard ~~Barometric~~ ^{Mercury} Temp. - 20°C
b) Height above Base & Mean Sea Level
c) Latitude - Standard is 45°
d) Instrumental Error.

- b) The Anoid Barometer - is less accurate than the Mercury Barometer but it is more compact and is self-recording. It consists of a metal box which expands and contracts with changes of pressure and may operate a pen giving a constant record of the pressure on a revolving drum. It is adapted for use as an Altimeter.

c) The Altimeter

An Altimeter is an Anoid Barometer which measures change of pressure and interprets it as change of height. Therefore we must assume some relationship between

pressure and height.

The Math in Altimeter assumes a constant temperature of 20°C and a Surface Pressure of 1013.2 millibars.

The Math in Altimeter assumes that the surface pressure is 1013.2 millibars and that the surface temperature is 15°C decreasing 1.98°C every 1000 feet up to 36,000.

An Altimeter therefore reads height above a fixed pressure and will only read height above a fixed level if the pressure at that level remains unaltered. If the pressure changes at the same level the Altimeter will indicate a change of height.

On the Math in Altimeter, the millibar subdial always gives the pressure at which the instrument would read 0 ft. This subdial can only be altered by hand.

If the air temperature does not agree with the standard conditions there will be a slight error in the indicated height. In warm air the Altimeter will underread.

Flying in warms air than standard alt. underreads.
" " colder air " " alt. overreads

5. Wind

On a weather map Isobars are formed joining places of the same Sea Level Pressure at a given time. They are drawn for every 2 millibars. The Wind at 2,000 feet blows parallel to these Isobars with Low Pressure on the left. (See Bay Banners Law) in the Northern Hemisphere. The wind speed is proportional to the closeness of the Isobars.

In the British Isles, if the Isobars are 6 miles apart, the wind is about 30 mph. If they are a hundred miles apart the wind speed is about 18 mph.

The closer the Isobars, the stronger the wind.

Surface Wind is generally estimated using the Beaufort Scale. This divides Winds into 12 ranges. Eg.

Force 0. - Winds of less than 1 mph.

Force 1. - Winds of from 1 to 3 mph.

Force 2. - Winds of from 3 to 10 mph.

Force 3. - Winds more than 10 mph.

With each Wind Speed a description of the

behaviour of common objects is given such as: Twigs, Paper, Leaves, Sea Surface etc. The direction is estimated by Smoke, Flags, Wind Socks, etc.

On a weather map wind is shown by an arrow with feathers to indicate the Beaufort Force.

Surface Winds may be measured with a cup anemometer incorporating an electric buzzer to count the revolutions. It does not record the direction but only the speed.

The Anemograph is a pressure tube which gives a constant ~~rec~~ record of the Wind speed on a rotating chart. A second pen gives the direction of the wind on another revolving chart.

- 11.0. When flying at about 2,000ft, Starboard Drift means that the Altimeter will over read 10. Flying from high to low pressure. If the wind is on the Port side, the m.s.l. pressure will decrease along the track. If the wind is 30 mph on the Port Beam, the Altimeter will overread 1 foot per mile.

Upper Winds are measured by Pilot Balloons and Pibalites.

The balloon filled with Hydrogen rises at the rate of 333 feet per minute and its bearing and elevation are measured each minute.

The Radio Balloon carries a transmitting set and its position is fixed at regular intervals by 3 ground stations. This method can be used to get wind at 60,000 feet.

The Nephoscope This method is used to find wind speed by watching the clouds. The height of the cloud is found by approximately times 1.6 miles High and Altitude - 3 miles. It is a useful method but not accurate. This method can only be used when either or both of the above two clouds are in the sky. The prongs of the Nephoscope are moved into the cloud direction - that gives Wind direction. The cloud is then timed between the prongs and that gives wind speed. This method can only be used between heights of 15,000-25,000 feet.

Day Savers are when standing with your back to the wind in the Northern Hemisphere, the low pressure is on your left. In the Southern Hemisphere the facts are reversed.

Friction

This is the variation in wind speed at the earth's surface produced by obstacles on the ground and surface friction.

In gusty winds, low level flying is often dangerous over buildings, hangars, etc. and great care should be taken.

Changes of Wind Direction

When wind alters its direction clockwise it is said to VEER.

When wind alters its direction anticlockwise it is said to BACK.

Diurnal Variation of Winds

Diurnal means, what happens throughout 24 hours.

- a. By Day, Surface Air Turbulence is marked, due to the earth becoming heated and the air above becomes warm, - expands and rises causing a mixing up of air.

The 1000 foot wind barometer and bulb is given up some of its force to the lower winds, which increase in speed and also tend to VEER. The surface winds usually attain their highest speed about the early afternoon.

7.7.0

3. At Night

The turbulence dies down and surface winds back and lull. The 1,500 foot wind freshen and V.S.R.

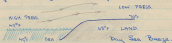
- The upper affected limit varies with, Phase, Season and Wind Strength.

Local Winds

1. Land and Sea Breezes

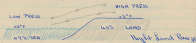
Daytime. The land becomes hotter than the sea, so the air above the land is heated and rises, while cooler air from over the sea comes in to take its place. It can be felt inland to about 10 to 15 miles and in the British Isles the speed is usually 10 to 15 m.p.h.

It is a normal feature of a quiet summer day. At 500ft. it is weak and it fades out altogether at 1000ft. Near the Equator it is stronger. Usually takes place from 1400 to 1700 hrs.



Night Time

The land cools more rapidly than the sea so the air over the sea rises and cools and heavier air comes to take its place from the land, causing a night land breeze which persists quietly until the morning unless upset by stronger disturbances. Its greatest effect is usually from 0100 - 0500 hrs.

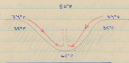


2. Katabatic Winds

Night time.

When air over high ground is colder than the air in the valley or on plains below, it tends to flow down the slope under gravity. This causes Katabatic Wind. The maximum effect is from 0100 to 0500 hrs. The slope determines the wind direction. All hills and mountains provide mild Katabatic Winds in fine quiet weather, more marked in winter time when snow is on the hills. It has a great effect

in causing radiation fog.

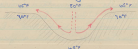


3. Mountain Winds

Day Winds

The sun heats the hillside and the air gets warm and rises and the air is hotter than the air at the same level over the valley.

Therefore the cold air sinks in the valley and tends to creep up the sides of the valley to take the place of the rising air. The effects are not so marked because the heavy air has to work against gravity. Therefore mountain winds are not as strong as katabatic winds. They take place from 15-1900 hrs.



6. Clouds

a) Humidity

The amount of Water Vapour that a volume of air can contain. Humidity is expressed in one or the other of the following ways:-

1) Relative Humidity - When the air is saturated it is known by the fraction.

- Weight of Water Vapour first present

Weight of Water Vapour when dew forms at the same temperature. It is expressed as relative humidity in % age.

Saturated air has a humidity of 100%

Day air 5%

2) The year average is about 70%

Dew Point

So the temperature at which the Water Vapour present in the air is sufficient to saturate it. When the temperature of the air falls below Dew Point, then water particles form - gives Dew. It is possible for the Dew Point to be less than the freezing point, therefore when the temperature falls below Freezing Point then Ice particles form giving Frost.

b) The Wet and Dry Thermometer

The dew point is found by using the Wet and Dry Thermometer.

This consists of 2 thermometers. The Dry one is a normal thermometer and the Wet one has a muslin bag round the bulb. Attached to the muslin bag is a Wick the end of which is in a small bowl of water. This keeps the muslin bag wet, - the water evaporates and therefore the air cools. This is shown on the Wet Thermometer. The difference in readings between the Wet and Dry Thermometer can be applied to tables from which the Dew Point can be determined.

c) Formation of cloud - stages.

1. A condition of rising air
2. Expansion owing to reduction in pressure and cooling
3. Less Water Vapour to saturate it
4. Cooling below Dew Point gives condensation into visible water drops.
5. Water drops are supported by up currents.

1. Convection Cloud - Cumulus or Cumulonimbus.

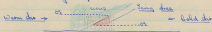
The heat of the Sun will heat the ground by day and the air in contact with the ground becomes warm and rises. If this is cooled to below its D Point then cloud will form. Since the warm rising air is replaced by cooler air, convection is momentarily stopped and so the cloud seldom covers the whole sky. Its base is usually between 500 to 5000 ft. and it shows marked vertical development with tops up to 50,000 feet in favourable conditions. These clouds disappear instantly at night when surface heating ceases. Over the sea they are formed by air movement and will tend to remain at night owing to the more variation of sea temperature.

2. Instability Cloud - Strato-cumulus.

Clouds, which are set up near the surface by obstacles and by friction, produce irregular up and down currents. If these are sufficiently vigorous and the air damp enough a layer of cloud is formed. Usually these clouds have quite a low base, generally between 1000-3000 feet thick.

3. Cloud of Warm Air over Cold. (Warm Fronted)

When a large mass of cold air meets a large mass of warm air then the warm air rises over the cold. As the warm air rises so it cools and condensation takes place and cloud is formed. Cloud formed in this way is usually of very wide extent both horizontally and vertically. A large sheet of cloud is formed extending from a very low level up to a height of about 5 miles.



4. Orographic cloud - Cumulus

This cloud is formed by air rising over high ground such as hills etc. The air does not follow the hill down but sweeps down gradually, therefore the cloud sinks and breaks up and disappears on the leeward side of the hill or high ground.



Thunder clouds are bulky in appearance and show marked vertical development. They are usually associated with showers since they are detached, and very large ones with Thunder storms.

Flying conditions are exceedingly bulky.

Layer clouds are usually formed by subsidence or at a Warm Front. They lie in nearly uniform sheets covering a wide area and give continuous drizzle or rain.

Cloud Classification

HIGH.	MEDIUM.	LOW.	Clouds of Great Vertical Extent.
20,000-50,000ft	2,000-20,000ft	Surface-2,000ft	10,000 - 50,000 feet.
Cirrus	Alto Stratus	Stratus	Cumulus
Cirrocumulus	Alto cumulus	Strato cumulus	Cumulo nimbus.
		Nimbo stratus	

(d) TYPES of

<u>Name</u>	<u>Location</u>	<u>Appearance</u>
Brown (C.)	South of Wam dis over hole	Siliceous ridge of about 100 crystals thick lying together and massive.
Brown streaks (C.)	South of Wam dis over hole	Thin whitish veil often covering hole they giving hole some sun - moon.
Brown mounds. (C.)	Instabilities at high levels	Large + patches of very small granular masses in waves + ripples. (called mounds)
Alto streaks. (A.)	South of Wam dis over hole	Gray veil covering the hole sky. Sun + Moon in but not through this.
Alto mounds. (A.)	Instabilities at medium levels	Large + patches of rather flattened masses, in lines waves + mounds.

Clouds

<u>Weather</u>	<u>Flying Conditions and A/c Values</u>
Approach of Wam front, indicated about 500 miles away when C. increases. Fine + threatening weather.	Too high + thin to affect normal conditions. Danger of condensation mists when flying at hi height.
Approach of Wam front indicated.	Cloud too thin to afford cover and too thin for normal use.
Fine	Nil.
Next approach of Wam front + transitional occasional slight drizzle.	Slight + mod. cumulus. Alto rising between 500 - 700. No aloft 1000 yds in this, + 1000 yds in thick fit. also instabilities. Affords not cover over wide area. 10-15000 ft thick.
Normally fine. banded type often indic's approach of front.	Slight cumulus. rising unsteady No cloud. cover afforded for A/c but useful for cohesion and distinction of course.

P.T.C.

Name	Location	Type of appearance
Stratus (Sg)	Turbidum or Orographical	Uniform layer resembling fog when at very low levels.

Stratus cumulus (Sc)	Rolls at low levels + spreading of convective cloud.	Extensive layer of globular masses in rolls, lines or waves often more shadowy
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Nimbo stratus (Ns.)	descent of warm low rise cold.	Does layer nearly uniform + if not with rain. More than 10,000 ft. patches of dark cloud above.
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Altostratus (As)	Convective	Thick cloud of Vert level at flat dark base - white cauliflower top. 20-30,000 ft. thick
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Clouds (Cont.) Weather	Flying Conditions and Ops Value
Occasional drizzle	Slight bumpiness, flying particularly in winter 0°C to -10°C. Low skillful + low to be of get information except as cast from ground observation. Dangerous, often covers high ground.

Normally dry, if any precipitation drizzle.	Most bumpiness in + below cloud - small above. Flying 0°C to -10°C but rate of descent not usually rapid. No 15-20 yrs often increase of time above cloud. Difficult not used + usually thin - 500-2000 ft thick
--	--

Continuous heavy rain, snow, sleet or drizzle.	Slight + most bumpiness in cloud. Severe in Turb. cloud below. Slight or med. flying. 0°C to -10°C. Severe below cloud of dark below freezing but Vis less than 20 yds Complete cover for 40% but use limited by flying conditions and low cloud base.
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Thunderstorm in cloud depth + height of up current.	Very bumpy in + below clouds. Heavy + rapid ascension. -10 to -15°C Vis less than 20 yds. No real sense useful only for ascension.
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P.T.O.

Types of

<u>Name</u>	<u>Location</u>	<u>Appearance</u>
Cumulus nimbus (C.N.)	Heavy inclement strong up-current.	Heavy mass of cloud. ft. Vert. development. Lower base, often flat having a fibrous texture + spreading out - cloud is deep. Thunder

Condensation Trails

At high altitudes, water vapors from the exhaust may condense giving a trail of cloud which spreads out and resembles times a Cirro cumulus. Dense trails rarely form below 28,000 ft. in summer and 22,000 ft. in winter.

Cloud cover - is rated and reported by the number of tenths of sky covered by each cloud type.

Measurement and Estimation of Cloud Height and Thickness.

By Day 1) Measurement. (a) Pilot Balloons.
(b) Special Met. Sights.
2) Estimation (b) Comparison with known Heights.

Clouds and

Weather

Heavy showers +
thunderstorms with hail
+ squally wind.

Flying Conditions + Air Value

Dense cloudiness in + below clouds.
- icing between 28-35°F. Also less
than 10 gals. danger from lightning.
No cont. cover. Only useful as a
last resort + then only for warning on
account of reverse flying conditions.

(b) Visual Estimation

By Night 1) Measurement. (a) Cloud Searchlight
(b) Met. Sight through cloud, to find thickness.

4. Visibility in Cloud

By visibility we mean the horizontal distance that can be seen on the ground with the naked.

By day it is measured by means of land miles whose distance from the MET office is known. The furthest object which can be clearly seen is the measure of visibility. Visibility may also be regarded as the transparency of the atmosphere, and at night, to measure this, fixed lights of known power and distance are used and viewed through a smoked glass of graduated shading. This is called the Visibility Meter.

The atmosphere may be observed in 2 ways.

- 1) Fog - made of Water Drops - is less than 1000 yds.
- 2) Mist - " " " " - Vis. between 1000 to 10000 yds.
- 3) Haze - made up of small particles of dust etc.
Vis. as mist

Haze is not so less confined to industrial areas. The sun side of deserts are fit for the effect from pollution of the atmosphere by chimneys.

Fog - being composed of Water Drops is produced by condensation at surface levels without any vertical movement of air.

There are 2 main types of Fog.

1) Radiation & Night Cooling Land Fog

This forms under the following conditions.

a) Clear Skies - Necessary so that earth may radiate its heat rapidly and so become cooled appreciably. If sky is cloudy the heat lost by the ground is re-radiated from the cloud layer and so no extensive cooling of the ground takes place.

b) Light Winds - less than 8 mph. Allow a shallow layer of cool air to form near the surface of the earth. The slight air movement is sufficient to produce mixing through this shallow layer and throughout this layer cooling takes place.

In a strong wind the turbulence effect takes place through too deep a layer and so clouds are more likely to form.

(c) High Relative Humidity.

(d) Cooling Below the Dew Point - Time factor.

Radiation Fog forms chiefly over low lying country and northland and is common in valleys owing to the drainage of cold air down the hillside at night.

Such fogs are most frequently found in the Autumn and Winter when the night radiation period is longest and usually forms in the early hours of the morning. Normally the fog clears later in the morning owing to heating by sunshine, but in Wintertime when the heat of the Sun is insufficient, the fog tends to persist until an increase in wind produces more turbulence and mixes the fog with warmer air above it causing it to evaporate.

2) Advection or Sea Fog.

This fog is formed when warm air moves over a cold surface. Such conditions usually being fulfilled at sea, though in very cold weather a sudden influx of warm air over frost bound ground may produce a fog.

Since the fog does not depend upon Radiation for its formation, it may form

under cloudy skies and in strong winds i.e. in the movement of the air.

When by contact the air at the surface is cooled to below its Dew Point - fog will form. Such fog may drift many miles inland and is most common in late spring and early summer when the sea is coldest.

Owing to the nonvariation of the sea temperature this fog may form at any time of the Day or Night.

Fog Thickness

The depth of Radiation fog increases with time. After 1 night, it is seldom more than 100 feet thick, but after several nights it may be in the order of 1,000 feet thick.

Sea fogs are of variable depths.

Fogs in Industrial Areas are always more dense since they are polluted with smoke and dust which cannot be carried away.

Precipitation - that which falls from clouds

The 4 main types of precipitation.

- 1) Rain.
- 2) Snow
- 3) Sleet
- 4) Hail.

1) Rain - The water droplets in a cloud which are supported by up currents gradually grow in size by further condensation and collision until they are too heavy to be supported by the up currents whence they fall to the ground as rain. The size of the droplets depends upon the strength of the up currents and the depth of the cloud. That is why the rain falls from heapier clouds is normally heavier than from layer.

2) Snow - When condensation takes place at a temperature below freezing point, ice crystals are formed. These grow in size by further condensation until they are too large to be supported and then provided that the temperature of the air

from the cloud base to the ground is below freezing point everywhere, snow will fall.

3) Sleet

Occurs when snow falls through a temperature a few degrees ($0^{\circ}\text{C} \pm 3^{\circ}\text{C}$) above freezing point and consists of a mixture of snow flakes (partly melted) and rain.

4) Hail - consists of hard pellets of ice produced by rain droplets being caught by swift up currents of air and carried up above freezing level long enough for them to freeze. In their subsequent movement above the freezing level in the cloud they grow by collision with super cooled water droplets which freeze immediately on impact, ultimately the hail stone falls to the ground, either because it is too heavy to be supported by the up current or because the up currents cease. Hail can be formed only in cumulo nimbus clouds.

9. Ice accretion

Ice forms on an aircraft when it encounters water drops at a temperature below freezing point. A constant watch must be kept on the aircraft thermometer particularly when entering clouds, remembering that these thermometers do not record the True Temperature unless a level flight for at least 2 minutes has been made.

Ice forms a danger owing to both increased weight and to change in wing section shape.

Types of Icing

a) Glassed Ice Accretion

The most important and serious type of icing. It forms all over an aircraft which is flying through air at a temperature below freezing point through which rain is falling. The rain must fall from a region where the temperature is above freezing point and this rain takes a long time to freeze as it falls, but on striking a solid object it will freeze immediately. The type of icing usually occurs below Nimbostratus at a Warm Front.

b) Ice Accretion is another serious type of icing and one which is more frequent and occurs when an aircraft flies through a cloud of water called water drops. These conditions are found in all water clouds between temperatures of 0°C to -15°C . These drops do not freeze into ice unless they are disturbed, but the aircraft producing that disturbance causes them to solidify on impact with it forming a transparent coating of ice.

Clouds in which Ice accretion is encountered.

i) Severe.

- | | |
|------------------|--|
| a) Cumulo nimbus | } 0°C to -10°C |
| b) Cumulus | |
| c) Nimbo stratus | -0°C to -7°C |

ii) Moderate.

- | | |
|------------------------------|---|
| a) Alto stratus | } 0°C to -4°C |
| b) Stratocumulus (not rapid) | |

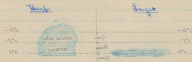
iii) Slight

- a) Stratus.

The severity of icing depends upon the size and number of super cooled water drops which in

them depend upon the strength of the up-currents. The size of the Water Drops decreases as the height increases and is gradually replaced by ice crystals.

In small clouds, the most severe icing occurs between 0° and -10°, and to a lesser degree down to -15°C.



Thunderstorms are the result of cumulo nimbus clouds which generate large electric charges and are ultimately able to break down the resistance of the air beneath the cloud. The spark which takes place is what we know as lightning and Thunder is the noise produced by it. For flying conditions and weather see the rainbow cloud. Compass and Radio are affected by the electrical discharge but danger from lightning is much serious unless the plane has no trailing wires.

10. The RAF Meteorological Service

The observations and information for preparing forecasts are distributed by the Central Office to the main types of stations.

They are:-

Type 1 At Headquarters of Bomber Command and Fighter Command and the Group Headquarters of these Commands. Their staffs are able to provide forecasts at once and at all times. They supply the forecasts and administer the system to 3 or 4 offices at headquarters of the Group and the offices of other Groups such as Training Command which have not a Type 1 office at Group.

Type 2 Found at bases such as operational stations of Coastal Command which require staff able to provide forecasts at any time. Their requirements cannot be fulfilled by a Type 1.

Type 3 These are attached to bases of Bomber Commands, C.M. of Coastal Command, A.C.M. and Ferry Pilot Pools. They obtain forecasts from their parent Group office (Type 1), but their staffs are qualified to brief cases and see.

captain Forewell.

Type II Found at Operational Control Room of Fighter Command where the staff obtain all information from the Type I office. Offices at stations which have night fighter squadrons have forecasters on duty all night.

A Weather Report - describes the weather at a given place at a given time.

A Weather Forecast outlines what weather conditions may be expected during flight. When a flight is to be made, a route forecast must be obtained on a Form 2330, and in requesting such a form the following information should be given, i.e. Either place or area + direction of flight, or Route and direction of flight.

Information in Standard Forecasts (Form 2330)

- 1) Date, Route & day, time and time of origin.
- 2) Brief account of the latest weather situation
- 3) The Weather.
- 4) Cloud (Type and Amount given in 10ths)

If high ground is likely to be in cloud then this is stated explicitly.

- 5) Freezing level given to nearest 500ft Mean SL. Estimated icing conditions in cloud is given as Severe, Moderate or Low.
- 6) If glazed frost is expected it is specially mentioned.
- 7) Visibility - given as a distance.
- 8) Surface Winds in Compass Points / MPH.
- 9) Upper Winds - Barzen / MPH.
- 10) Barometric Pressure at Home Base (QFE) of destination airfield and likely diversion airfields and at MSL (QFF) for places en Route.
- 11) The Sun & Moon Rising and setting times and phases of moon.

NOTE On operational flights, place names are omitted and Item 2. is secret.

The Form 2330 is completed by the members of the aircraft crew and returned to the MET. office after the flight.

Observations en Route are made and the pilot is to note alt., time, Position, Altimeter Readings, Cloud Type, Amount, height of base and top if possible, Temp. and Visibility.

Under Route Forecasts, the following are issued as required:-

1. Gale Warnings - issued when surface wind in the open is expected to exceed 14.0 mph. The max force of gale is usually given.
2. Squalls or Thunderstorms - if after approach the stations a warning is issued at once.
3. Frost Warning - issued when the air temp is likely to fall below 32°F. In winter when a frost is a daily feature of the weather, a warning is only issued only when the frost is likely to become more severe. The minimum temp is always given.
4. Snow Warnings - issued so that preparation may be made to keep runways clear.
5. Local Forecasts - issued 3 or 4 times a day to cover a period of about 8 hrs and a radius of 20 miles.

Basis of Forecasting

An air mass is a large volume of air whose characteristics of temperature and humidity are constant throughout the mass. As the air mass moves, it will bring with it the characteristics of its latitude or region of origin.

As it drifts to other latitudes, from what we know of temperature, pressure, and humidity, these will give rise to weather changes.

A. Factors affecting the Weather

1. Air mass
2. Pressure distribution and fronts
3. Climate and Seasons, Shape, Size and position of Land and Sea.

B. Polar Air Weather

Air, cool and dry, warmed at the surface causes convection cloud. Showers, perhaps Thunder, gusty Surface Winds, Good Vis, Bumpy flying by day but smoother at night, and icing in cloud.

Tropical Air Weather

Air warm and moist, cooled at the Surface - little or no convection, layers of cloud formed, usually Stratus, Drizzle, Poor Vis. and fog. Slight or moderate bumpiness, slight icing.

11.

FRONTS

A front is a zone of transition between two different air masses, and on a weather map are drawn as a line.

a. Weather Change with a WARM front.

If, after a spell of cold weather with good visibility, the clouds become gradually thick and low with a period of continuous rain and drizzle, the wind veers and the temperature rises and the weather is warm, muggy, drizzly and overcast with poor visibility, then a warm front has gone through.

b. Weather Change with a COLD front.

If, after a spell of warm, muggy, drizzly overcast weather, the clouds become very thick with very heavy rain, the wind veers and the temperature falls, followed by a

decrease and then bright intervals and showers with good visibility, then a COLD FRONT has gone through. (See diagram at foot of page).

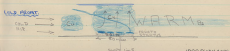
2. a. Beaufort Weather Notation.

22 Not to be confused with the Beaufort wind scale.

Originated as a uniscribed method for noting down weather observations in a concise form. Letters were first used but these have been gradually replaced by a figured code, though the letters are still used to a slight extent.

b. Information on a Weather Map.

- 1) Pressure reduced to M.S.L.
- 2) Pressure change in last 3 hours.
- 3) Temperature.
- 4) Dew Point.
- 5) Visibility.
- 6) Wind Force and Direction.
- 7) Clouds, Type and Height of Base.
- 8) The Weather.



D.R. Navigation

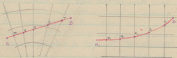
INDEX

1. Plotting Gt. Circle Brg on a Mercator Chart, using a Gnomonic Chart.
2. Drift and Wind Same Method of Wind Finding.
3. The Squat Search.
4. Position Lines: a) Methods of Obtaining Pos. Lines. b) Use of Single Pos. Lines.
5. i) Drift. a) Importance of. b) Methods of Obtaining. ii) Methods of Obtaining T.M.G.; b) Measuring Gt.
6. Mean Equivalent W/V (for climbing).
7. Wind Finding. Relative Advantages and Disadvantages.
8. Navigational Log. - Form 1941.
9. Effects of Changes of W/V on Drift and Gt.
10. Methods of Fixing Position Lines.
11. Critical Point - between Two Bases (also see after 12).
12. Revised Method of Plotting Position Lines. Curve of Equal Brg.
13. Reasons for Deviation from most Direct Course.
14. Flying at Various Heights - Advantages and Disadvantages.

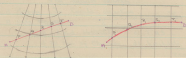
1. How to plot a Gt. Circle Brg on a Mercator Chart using a Gnomonic Chart.

- a) Join the 2 places concerned on the Gnomonic Chart.
- b) Note the latitudes where the track cuts various meridians. (As in diagram P.Q.R.S.T.)
- c) Transfer these positions by Lat. and Long. to a Mercator Chart (P, Q, R, S, T.)
- d) Join these positions freehand.

The resultant curved lines represents a Gt. Circle on a Mercator Chart.



Southern Hemisphere



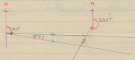
Northern Hemisphere

2. Drift and Wind Lane Method of Wind Finding.

Depends on the fact that the wind veers with height by 10° for 1,000 feet.

Example. Co. 100°T, T.A.S. 120; Height 1500 ft.
Method.

- 1) Measure Drift. (Say 5°)
- 2) Measure the ground direction of the wind lanes by means of Bearing & Altitude Compass (Say 058° & 215°)
- 3) Since there is Port Drift, the Surface Wind direction is from 215°
- 4) Since Wind veers with height, wind at 1,000 ft will be from 225°
- 5) Therefore the Triangle of Velocities may be solved.



Disadvantages

- 1) Cannot be used at night, though Wind lanes can be seen by shining the Searchlight of the ship on them. This is impracticable during wartime.
- 2) Can only be used at heights of 1,000 ft or less.
- 3) Can only be used over sea.
- 4) There must be sufficient Surface Wind to form Wind Lanes.
- 5) Cannot be used for flying up & down wind, i.e. you must be making a Drift.

Advantages

- 1) No alteration of course necessary.
- 2) Quick and accurate.
- 3) Simple unnecessary.
- 4) It is minute Wind Velocity.
- 5) Easily provides a check on Ground Speed.

2.

The Square Search

1) Returns To Locate objective

2) Route Required

- | | |
|--------------------------|----------------|
| a) Wind Velocity. | c) T.A.S. |
| b) DR Pair of Objective. | d) Visibility. |

P.S.O.

3. Importance of Observation

- a) By off up or down Wind to make good a track equal to 2V (True Variability)
 - b) Lay off track at 90° equal to 2V
 - c) " " " " 90° " " 4V
 - d) " " " " 90° " " 14V
 - e) " " " " 90° " " 6V
- etc.

4. The possibility of Error

1 only - Underestimate V rather than overestimate it.

5. Practical Application

- a) Work out all courses and ETA's for all tracks in search before arriving at the DR position.
- b) Failing to locate objectives, alter the following course (see tables)
- c) Check W/P at turning point by Double Drift
- d) Plot courses and air distances from the air position from which commence search. AD Plot Master from last fix.

Course	Dist	Geo	Geo	Time	Dist	DATA
164	12	165	164	14	14	W/P = 119/054
214	12	213	213	34	13	TAC = 2204
204	20	205	207	54	19	Via = 64mm.
094	20	213	091	64	25	

ca

1.

Position Lines

a) Methods of Obtaining Position Lines

- 1) By observing 2 known object in transit
- 2) By bearing compass
- 3) By bearing place or helio compass (Rel. Brgs.)
- 4) By Loop. (Rel. Brg.)
- 5) W/P D/P.
- 6) Astronomical Observations
- 7) By determining Range of a known object.

b) Uses of a Single Position Line

1. To check G.B.

Eg. Track 090, fix a weather
rain lines at 1030 hrs. from
1010

If on track, then will have flown distance
4.5 x in 20 mins.

NB. To check G.B. position lines should not cut the
track at an acute angle.

2. To confirm the track

Track 090, Obtain rain line
i.e. Not far off track.

NB. Must make a small angle with the track.



3. To give the Regid Incho to a Certain Position

- a) By Visual Browsing
- b) By W.T. M.F. Browsing
- c) By Deep Browsing



If a bearing can be taken of B then the bearing can be the back to B.

5 4 Daily

- a) The Importance of Drift
 - 1) Determine the T.M.S.
 - 2) It acts as a check on W/P.
 - 3) Provides a means of determining W/P.
 - b) Methods of Obtaining Drift
 - 1) C&BS.
 - 2) Drift Recorder
 - 3) Rest Street. Gun trained on some object that the ship has recently passed over.
 - 4) Back Bearing.
 - 5) By determining T.M.S.
- 2) a) Methods of Obtaining T.M.S.
- 1) By 2 fixes
 - 2) By measuring Drift and applying to G.C.T.

- a) By moving a Bush Bearing.
- b) By V.T. Bp Bearings (In certain cases, i.e. bearing pressed over the shaft and not allowed to.)
- c) By GDS's and GDS's.
- d) By Lock Bearings (bearing fixed over shaft)
- e) By 3 or more position lines from one or more engines.

b) Measuring Growth Speed

- 1) 1882.
- 2) By Drift Recorder
- 3) By 2 fms
- 4) By position lines cutting the track almost at right angles

b. Mean Equivalent Life (for Blinding)

The method used when using 2, 3 or 4 words, but the right interval between the given mix is $\frac{1}{2}$ must be equal and the Rule of 4 must be steady.

1. Max Wind at 5,000 ft: $140/30 = 4.6$
2. " " " 10,000 ft: $200/40 = 5.0$

and 94.6% of the total variance.

- $$\log m = \log 2 + \log n \quad \text{on table of fig. 5.}$$

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Figure 6

Vol. 1: Issues 1-4 for 2000: 14, 15, 16, 17, 18, 19, 20, 21, 22, 23, 24, 25, 26, 27, 28, 29, 30, 31, 32, 33, 34, 35, 36, 37, 38, 39, 40, 41, 42, 43, 44, 45, 46, 47, 48, 49, 50, 51, 52, 53, 54, 55, 56, 57, 58, 59, 60, 61, 62, 63, 64, 65, 66, 67, 68, 69, 70, 71, 72, 73, 74, 75, 76, 77, 78, 79, 80, 81, 82, 83, 84, 85, 86, 87, 88, 89, 90, 91, 92, 93, 94, 95, 96, 97, 98, 99, 100, 101, 102, 103, 104, 105, 106, 107, 108, 109, 110, 111, 112, 113, 114, 115, 116, 117, 118, 119, 120, 121, 122, 123, 124, 125, 126, 127, 128, 129, 130, 131, 132, 133, 134, 135, 136, 137, 138, 139, 140, 141, 142, 143, 144, 145, 146, 147, 148, 149, 150, 151, 152, 153, 154, 155, 156, 157, 158, 159, 160, 161, 162, 163, 164, 165, 166, 167, 168, 169, 170, 171, 172, 173, 174, 175, 176, 177, 178, 179, 180, 181, 182, 183, 184, 185, 186, 187, 188, 189, 190, 191, 192, 193, 194, 195, 196, 197, 198, 199, 200, 201, 202, 203, 204, 205, 206, 207, 208, 209, 210, 211, 212, 213, 214, 215, 216, 217, 218, 219, 220, 221, 222, 223, 224, 225, 226, 227, 228, 229, 230, 231, 232, 233, 234, 235, 236, 237, 238, 239, 240, 241, 242, 243, 244, 245, 246, 247, 248, 249, 250, 251, 252, 253, 254, 255, 256, 257, 258, 259, 260, 261, 262, 263, 264, 265, 266, 267, 268, 269, 270, 271, 272, 273, 274, 275, 276, 277, 278, 279, 280, 281, 282, 283, 284, 285, 286, 287, 288, 289, 290, 291, 292, 293, 294, 295, 296, 297, 298, 299, 300, 301, 302, 303, 304, 305, 306, 307, 308, 309, 310, 311, 312, 313, 314, 315, 316, 317, 318, 319, 320, 321, 322, 323, 324, 325, 326, 327, 328, 329, 330, 331, 332, 333, 334, 335, 336, 337, 338, 339, 340, 341, 342, 343, 344, 345, 346, 347, 348, 349, 350, 351, 352, 353, 354, 355, 356, 357, 358, 359, 360, 361, 362, 363, 364, 365, 366, 367, 368, 369, 370, 371, 372, 373, 374, 375, 376, 377, 378, 379, 380, 381, 382, 383, 384, 385, 386, 387, 388, 389, 390, 391, 392, 393, 394, 395, 396, 397, 398, 399, 400, 401, 402, 403, 404, 405, 406, 407, 408, 409, 410, 411, 412, 413, 414, 415, 416, 417, 418, 419, 420, 421, 422, 423, 424, 425, 426, 427, 428, 429, 430, 431, 432, 433, 434, 435, 436, 437, 438, 439, 440, 441, 442, 443, 444, 445, 446, 447, 448, 449, 450, 451, 452, 453, 454, 455, 456, 457, 458, 459, 460, 461, 462, 463, 464, 465, 466, 467, 468, 469, 470, 471, 472, 473, 474, 475, 476, 477, 478, 479, 480, 481, 482, 483, 484, 485, 486, 487, 488, 489, 490, 491, 492, 493, 494, 495, 496, 497, 498, 499, 500, 501, 502, 503, 504, 505, 506, 507, 508, 509, 510, 511, 512, 513, 514, 515, 516, 517, 518, 519, 520, 521, 522, 523, 524, 525, 526, 527, 528, 529, 530, 531, 532, 533, 534, 535, 536, 537, 538, 539, 540, 541, 542, 543, 544, 545, 546, 547, 548, 549, 550, 551, 552, 553, 554, 555, 556, 557, 558, 559, 560, 561, 562, 563, 564, 565, 566, 567, 568, 569, 570, 571, 572, 573, 574, 575, 576, 577, 578, 579, 580, 581, 582, 583, 584, 585, 586, 587, 588, 589, 590, 591, 592, 593, 594, 595, 596, 597, 598, 599, 600, 601, 602, 603, 604, 605, 606, 607, 608, 609, 610, 611, 612, 613, 614, 615, 616, 617, 618, 619, 620, 621, 622, 623, 624, 625, 626, 627, 628, 629, 630, 631, 632, 633, 634, 635, 636, 637, 638, 639, 640, 641, 642, 643, 644, 645, 646, 647, 648, 649, 650, 651, 652, 653, 654, 655, 656, 657, 658, 659, 660, 661, 662, 663, 664, 665, 666, 667, 668, 669, 670, 671, 672, 673, 674, 675, 676, 677, 678, 679, 680, 681, 682, 683, 684, 685, 686, 687, 688, 689, 690, 691, 692, 693, 694, 695, 696, 697, 698, 699, 700, 701, 702, 703, 704, 705, 706, 707, 708, 709, 710, 711, 712, 713, 714, 715, 716, 717, 718, 719, 720, 721, 722, 723, 724, 725, 726, 727, 728, 729, 730, 731, 732, 733, 734, 735, 736, 737, 738, 739, 740, 741, 742, 743, 744, 745, 746, 747, 748, 749, 750, 751, 752, 753, 754, 755, 756, 757, 758, 759, 760, 761, 762, 763, 764, 765, 766, 767, 768, 769, 770, 771, 772, 773, 774, 775, 776, 777, 778, 779, 780, 781, 782, 783, 784, 785, 786, 787, 788, 789, 790, 791, 792, 793, 794, 795, 796, 797, 798, 799, 800, 801, 802, 803, 804, 805, 806, 807, 808, 809, 810, 811, 812, 813, 814, 815, 816, 817, 818, 819, 820, 821, 822, 823, 824, 825, 826, 827, 828, 829, 830, 831, 832, 833, 834, 835, 836, 837, 838, 839, 840, 841, 842, 843, 844, 845, 84

The drawing: Some fantastic birds you can't find

Wind Finding - Relative

Method

1. Multiple Drift

Advantages

- Up to the Minute Wind (Bombing)
- Does Not require fix
- Provides its own check in accuracy
- Obtained over sea and land
- Little time lost, but also time used.
- Need not be plotted

2. Dir Plot

- Gives mean wind which may be very advantageous for return flight.
- Gives free freedom of action
- No alteration of course necessary.
- Several passing off points and fix may be accurately obtained.
- Promotes interest in checking course, time and time.

3. Track and G/L

- No loss of time - no alteration of Co. necessary
- Simple, c) No plotting required.

4. Drift + Wind Line

- W/F found out sea without fix.
- Alteration of Co. unnecessary
- No loss of time.
- Only 1 accurate drift required.
- No timing necessary.
- Need not require plotting. 200

Advantages and Disadvantages

Disadvantages

- Alteration of course necessary.
- Not always possible to obtain Drifts (Land)
- Certain amount of time lost
- Local Wind which may not have any bearing of future conduct.
- Plotting and courses must be accurate
- Reliable fixes must be obtained otherwise W/Fs found become merely check winds.
- The mean W/F may have no bearing on the actual W/F at a particular time (for use for Bombing).
- Geographical ground points are necessary.
- Several time necessary.
- Does not allow free freedom of action.
- Course and W/F must be kept steady.
- Can be found over sea only
- Good Visibility required.
- Depends upon the assumption that wind VEERS 10° with 1500 ft in height.
- Little practical value in excess of 1500 ft. 200

Drift and Wind Lane Method of Tracking Wind (cont) Advantages (continued)

- g) Gls obtained in same operation.
- h) Up to the minute wind.

Disadvantages (continued)

- a) Sea must not be too rough or too calm otherwise Wind Lane cannot be observed.
- f) Cannot be used when flying up or down wind.
- g) Cannot be used at night.
- h) Local Wind - may not have any bearing on future conduct.

8. Bom logs

The purpose of Bom keeping

- 1) To prevent mistakes being made in the air. Dangerous to commit to memory in the air, Bom gives a written record.
- 2) To enable another member of the crew to take over should the Navigator become a casualty.
- 3) To enable the flight to be constructed, so that navigators can justify any action taken.
- 4) Gives a record of actual weather experienced, also information useful to Intelligence.
- 5) Gives a record of engine and airframe times.

- 6) Give a record of life signals sent and messages received.
- 7) Can be used in court of law as evidence.

9. The Effects of Changes of W/P on Drift and Gls

1. Head Winds	<u>Effect on</u>	
	<u>Drift</u>	<u>Gls</u>
a) Change in Direction	Large	Little



2) Change in Speed	<u>Effect on</u>	
	<u>Drift</u>	<u>Gls</u>
	Little	Large



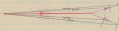
2. Fan Winds

a) Change in Direction



Effect on-	
Drift	GLS
Mod. Large	Little.

b) Change in Speed



Drift	
GLS	
Little	Large

3. Beam Winds

a) Change in Direction



Effect on-	
Drift	GLS
V. Small.	Large.

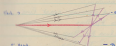
b) Change in Speed



Drift	
GLS	
Large	Small.

4. GUSTY Winds

a) Change in Direction



Effect on-	
Drift	GLS
Mod. Large.	Mod. Large.
Small.	Mod. Large.

b) Change in Speed



Drift	
GLS	
Mod. Large.	Mod. Large.
Large.	Mod. Large.

10. Methods of Fixing Position Lines. Summary in order of Accuracy.

1) Visual Fix

- A Visual Ground Position
- By 2 or more Visual position lines.

2) Astro Fix

- Fixes class Astro sights accurate plus or minus 10 miles irrespective of distance from any means of navigational aid. Improvement

in Sights will reduce error considerably in the future.

3. WV Lines in order of accuracy:-

- a) First class WV Brg and crossing coast line
- b) 2 or more first class WV Brgs
- c) 2 or more second class WV Brgs
- d) 2 or more third class WV Brgs
- e) Loop Brgs - cannot be assessed

- Note:-
- 1st class WV Bearing is accurate to $\pm 2^\circ$
 - 2nd " " " " " " " " $\pm 4^\circ - 5^\circ$
 - 3rd " " " " " " " " from 5° upwards
- 2° error will give an error of about 2 miles at a distance of about 60 miles. The accuracy of the fix will therefore depend upon:-
- a) Distance from the Station
 - b) Angles at which the lines cross each other, -
the more ^{acute} ~~acute~~, the greater the possibility of error.

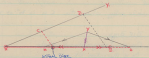
4. The Critical Point - between 2 fixed Brgs.

Definition - is the point from which it is equally quick to fly to either base.

- WV Brgs
- a) Reduced Air Speed. a) WV
 - b) Time. b) Distance.
 - c) Track.

2. Methods of Finding the Critical Point.

1) Graphic



- a) Draw Track A-B
- b) Choose any point along AB, as X and lay off WV, as
- c) Draw arcs (Radius = 200) from Y cutting Track at H+O
- d) Therefore as. fly back = 20. fly out.
- e) Lay off a line from A at any angle AY.
- f) Mark off fly back along AY - C. and then from C mark off fly out along AY - D.
- g) Join D and B.
- h) Join C to track, parallel to DB and where it cuts the track is the Critical Point.

2) Out and Home Formula

Time. ^{Dist} Speed	=	
X. fly Home	$\therefore \frac{Dx}{Sx} = \frac{Dx}{Hx}$	$\therefore Dx \cdot Hx = 20.$
O. fly Out.	$\therefore Dx = 20 \div Hx$	$\therefore Dx = 2(6.4)$
	$\therefore x = \frac{DxH}{SxSx}$	= Distance formula.

3. Refined Method of Plotting Position Lines Introducing the curve of course correction

Applicable only when the CA is 10° or more.

1) True D.R. Triangle

- a) On a Topographical Map
On a Mercator Chart

Method of laying off Position Line

- 1) Calculate D.R. position of 40° or, if impossible by lot assume a position.

- 2) Calculate CA

- 3) Apply CA to get True D.R. (Given by the station)
to determine a True Line Bearing

- 4) Lay off RL D.R. from station.

Since Wireless Waves follow the great circle track, it would be more accurate to assume that the Position Line is part of the great circle track about your D.R. position.

- 5) Transfer from position where True Line Bearing into your True D.R. position draw a line towards the station and away from the equator inclined to the True Line by the CA you have previously calculated.

4) Mercator Chart (cont.)

1. Northern Hemisphere



2. Southern Hemisphere



← Refined Pos. Line

1) On a Topographical Map

Method of laying off Position Line

- 1) Lay bearing off direct from the W.T. Station as a straight line.

1. Northern Hemisphere



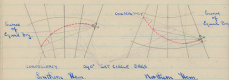
2. Southern Hemisphere



2. Loop Bearings

a) On a Topographical Map

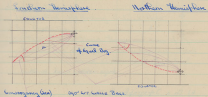
A bearing say 90° laid off from the M/C can be laid off to the station from every meridian. (see diagram). If all these points where each bearing crosses the respective meridian were joined together a curve concave to the Pole would be produced.



The Curve of Equal Bearing - is a line joining all places which have the same Great Circle Bearing of a given position at any instant. It constitutes a more accurate position line obtained from D/F Loop Bearings.

b) On a Mercator Chart

A Great Circle Bearing of 90° is laid off from the M/C to a station from every meridian. If all the points where the Great Circle Bg cut the respective meridians were joined, the result would be a curve concave to the Pole. (see diagrams).



Method of Plotting the Tangent of the Curve of Equal Bg

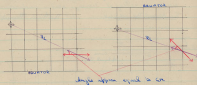
- 1) Work out Loop Bg in the normal way and Plot the Rhumb line Bearing from the Station
- 2) Where the Rhumb line Bg cuts the Ed. Meridian draw a line inclined at an angle of CA to the R. to connect the 1° Station and towards the Equator. See

This will represent a tangent to the curve of Equal Bearing and is a more accurate position line to use from Equal Bearings. (See Diagram).

a) On a Mercator Chart

Northern Hemisphere

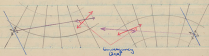
Southern Hemisphere



b) On a Spheroidal Map

Northern Hemisphere

Southern Hemisphere



Plotting the Curve of Equal Bearing on a Spheroidal Map

This is carried out the same way as on a Mercator Chart except that your D.R. meridian must be transferred to pass through the D.R. Pt. This is necessary because the True Brg was taken relative to your D.R. meridian.

NO. The True Great Circle Bearing is plotted, not the Rhumb Line Bearing.

Amendment

2nd Method of finding Critical Point

ie By Computer

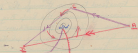
Place Ground Speed Down over sum of Ground Speeds Out and Home and against Total Distance on inner scale, read off critical distance on outer scale.

Reasons for Deviation from most direct route.

- 1) To make sure of visual fixes.
- 2) To avoid well defended areas.
- 3) To avoid unfavourable weather (See diagram).
- 4) To avoid difficult country i.e. jungle, Desert and long stretches of sea.
- 5) Practical reasons - to give surprise.

c) To be in accordance with 53.133 (Prohibited Areas)

(3) To avoid unfavourable weather.



Cold Front Warm Front

a) To fly along direct route from A to B results in contrary winds, Heavy Rain and perhaps icing from Warm Front, and Heavy rain perhaps icing with possible Thunderstorms in Cold Front.

b) To fly along Indirect Route from A to B you avoid the Fronts and will have favourable winds.

2. Flying at Various Heights: Advantages and Disadvantages

1. Low Flying

a) ADVANTAGES

- 1) If visibility permits, frequent in-coming and backing of Drift a destination will be eventually be reached and/or a specified area reached.

b) DISADVANTAGES

- 1) In bad visibility, no check can be kept on G/S or W/P.
- 2) If pilot does worse without fixing the position of the A/C.
 - 1) A new ILL posn may not be found or
 - 2) A course to judge good & Required Track.

2. Low Flying

a) DISADVANTAGES (500 feet or less)

- 1) View restricted
- 2) Pass. unable to maintain steady G.S. & A/S speed
- 3) Time difficult to obtain ILL
- 4) Alt. observations impossible
- 5) ILL Range severely restricted (Sinking and cannot be ruled out)
- 6) Close up firing only possible

b) DISADVANTAGES

- 1) Safety
- 2) Surprise

3. Medium Altitude Flying (500 to 10,000 ft)

a) ADVANTAGES

- 1) Ground often visible - enables visual firing.
- 2) Drifts easily obtained.
- 3) W/P's checked with accuracy

- i) Good forecast of winds.
- j) Able to make use of cloud cover if necessary.
- k) Cold usually not too severe.
- l) No Oxygen necessary.

b) ~~Intermediate~~

- i) Radio observations may be restricted. Cloud.
- j) Range of D/F less than at great heights.
- k) Enemy opposition likely to be greatest.

c) High Flying (above 10,000 feet)

a) ~~Intermediate~~

- i) Radio observation conditions good.
- j) D/F facilities good with good range.

b) ~~Intermediate~~

- i) Use of Oxygen necessary.
- j) Cold can be very severe.
- k) D/F difficult to obtain, even if ground is visible.
- l) Visual enemy difficult.
- m) W/F's difficult to forecast and often very strong.
- n) Assets and finding difficult.